



Stable compounds in the CaO-Al₂O₃ system at high pressures

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Abstract

Using the evolutionary crystal structure prediction algorithm USPEX, we showed that at pressures of the Earth's lower mantle (24–136 GPa) CaAl₂O₄ is the only stable calcium aluminate. At pressures above 7.0 GPa it has the CaFe₂O₄-type structure with space group *Pnma*. This phase is one of the prime candidate aluminous phases in the lower mantle of the Earth. We show that at low pressures 5CaO • 3Al₂O₃ (C₅A₃) with space group *Cmc*2₁, CaAl₄O₇ (with space group *C2/c*) and CaAl₂O₄ (space group *P2₁/m*) are stable at pressures of up to 2.1, 1.8 and 7.0 GPa, respectively. The previously unknown structure of the orthorhombic "CA-III" phase is also found in our calculations. This phase is metastable at 0 K and has a layered structure with space group *P2₁2₁2*.

Keywords CaO-Al₂O₃ system · Crystal structure prediction · High pressure · Evolutionary algorithm · USPEX · Earth's mantle

Introduction

Aluminum is the fifth most abundant element in the Earth's crust and mantle, the Al/Si ratio being 0.36 in the crust and 0.17 in the mantle (Anderson 1983). Today, it is thought that the lower mantle consists mainly of (Mg,Fe)SiO₃ bridgmanite with a perovskite-type (Pv) crystal structure (~70 vol.%), (Mg,Fe)O ferropericline (~20 vol.%), and calcium perovskite CaSiO₃ (Ca-Pv) (6–12 vol.%) (Ringwood 1975; Irifune 1994; Pushcharovsky and Pushcharovsky 2010). It is known that in rocks of the composition of peridotite and pyrolite, Al is incorporated into bridgmanite and there is no separate aluminous phase at least up to the middle region of the Earth's lower mantle (Irifune et al., 1994; Kato et al. 1996; Liu et al. 2016). Several potential aluminous phases have been considered for the lower mantle. For example, a

solid solution between MgAl₂O₄ and CaAl₂O₄ is known (Akaogi et al. 1999). It was also previously shown that Al-rich basalts contain MgAl₂O₄ with the CaFe₂O₄-type structure, which could coexist with aluminous bridgmanite in the Earth's lower mantle (Litasov and Ohtani 2005). Thus, the crystal chemistry of the CaO-MgO-Al₂O₃ system plays an important role for understand the chemistry of Al in the Earth's lower mantle.

All experimentally known polymorphic modifications of CaAl₂O₄ can be divided into low-pressure structures, where which aluminum atoms are in the tetrahedral coordination, and high-pressure phases, with aluminum in five or sixfold coordination. The well-known monoclinic modification with the space group *P2₁/n* at normal conditions (Ito et al. 1980) consist of a framework of AlO₄ tetrahedra, topologically equivalent to the tridymite structure (SiO₂) (Fig. 1a). The voids of the framework are occupied by calcium atoms in three independent crystallographic sites with different coordination numbers: 6 (Ca1), 6 (Ca2) and 9 (Ca3). This modification was first observed in meteorites and the mineral was named crotite (Ma et al. 2011). Another monoclinic modification of CaAl₂O₄ with space group *P2₁/c* (Fig. 1b) crystallizes in the *m*-CaGa₂O₄ structure type (Ito et al. 1980), where the AlO₄ distorted tetrahedra form a three-dimensional framework with the tridymite topology, and Ca atoms are seven-coordinate. This phase was also found in meteorites and was named dmitriivanovite (Mikouchi et al.

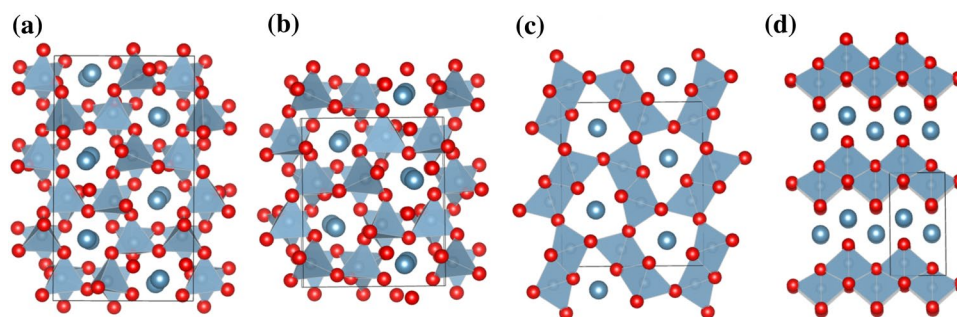
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Fig. 1 Polyhedral representation of experimentally known CaAl_2O_4 structures with AlO_4 tetrahedra (gray color): $P2_1/n$ (xz projection) **a**, $P2_1/c$ (yz projection) **b**; and AlO_6 octahedra (gray color): $Pnma$ modification (xy projection) **c** and "layered" $P2_1/m$ modification (zy projection) **d**. Ca^{2+} ions are shown as large balls



2009). The monoclinic $P2_1/c$ phase is slightly denser than $P2_1/n$, and the transition from $P2_1/n$ to $P2_1/c$ is observed at the pressure of 1 GPa and the temperature of 700 °C and at 2 GPa and 1300 °C (Ito et al. 1980). Among other low-pressure modifications of CaAl_2O_4 with AlO_4 tetrahedra, a crystal structure with space group $P6_3$ has been mentioned (Janáková et al. 2007). However, the authors have questioned the correctness of their structure refinement because of the high R-factor and highly distorted interatomic distances. Most likely, this phase is metastable at low pressures and does not have a stability field in the phase diagram.

With further increase in pressure, the coordination of Al atoms changes from tetrahedral to octahedral. Among the CaAl_2O_4 phases with the octahedral coordination of Al atoms the best-known orthorhombic $Pnma$ modification was first described in 1957 (Decker and Kasper 1957) (Fig. 1c). This phase crystallizes in the CaFe_2O_4 structure type (space group $Pnma$) where edge-sharing AlO_6 octahedra forming channels along the c axis, filled with Ca^{2+} cations.

Another modification of CaAl_2O_4 is the monoclinic $P2_1/m$ phase with a layered structure, synthesized by (Lazić et al. 2006) (Fig. 1d). The layers of AlO_6 octahedra in this structure are separated by interlayer Ca cations. The phase transition $P2_1/m \rightarrow Pnma$ has been experimentally discovered at $P = 4\text{--}8$ GPa (Lazić et al. 2006). In the experimental study of the $\text{CaAl}_2\text{O}_4\text{--CaGa}_2\text{O}_4$ system (Ito et al. 1980), yet another modification of CaAl_2O_4 , named "CA-III", and stable in the pressure range of 3–8 GPa, was described. Its crystal structure was not determined, but an orthorhombic cell was established with the parameters $a = 4.39$ Å, $b = 5.07$ Å, $c = 6.96$ Å.

In a recent theoretical paper (Eremin et al. 2016), it has been shown that the phase transition between $P2_1/m$ and $Pnma$ is predicted by *ab initio* calculations at 7–8 GPa, which is in close agreement with the experiment (Lazić et al. 2006). Calculations based on semiclassical interatomic potentials overestimate the phase transition pressure, 18–19 GPa (Eremin et al. 2016). As calculations (Eremin et al. 2016) showed, when pressure increases, $Pnma\text{--CaAl}_2\text{O}_4$ becomes increasingly more favorable with increasing pressure in the entire range of mantle pressures

and temperatures. The $Cmcm$ phase of CaAl_2O_4 is always slightly (0.2–0.5 eV/f.u.) less stable than the $Pnma$ phase.

Analyzing recent experimental and theoretical studies, one can conclude that under $P\text{--}T$ conditions of the Earth's mantle, only two phases of CaAl_2O_4 may exist: monoclinic $P2_1/m$ and orthorhombic $Pnma$, both with the octahedral coordination of the Al atoms, and the stability field of the latter covers almost the entire $P\text{--}T$ range of the Earth's mantle. So far, only the phases of stoichiometry CaAl_2O_4 have been investigated at conditions of the Earth's mantle, and even there some structures (such as Ca-III) remain experimentally unresolved.

In addition to the structural modifications of CaAl_2O_4 , literature describes experimentally investigated phases with different stoichiometric ratios where the aluminum atoms have tetrahedral coordination ($\text{Ca}_{12}\text{Al}_{14}\text{O}_{33}$ (Büsemann and Eitel 1936), CaAl_4O_7 (Boyko and Wisnyl 1958), $\text{Ca}_3\text{Al}_2\text{O}_6$ (Mondal and Jeffery 1975), $\text{Ca}_5\text{Al}_6\text{O}_{14}$ (Vincent and Jeffery 1978), $\text{Ca}_4\text{Al}_6\text{O}_{13}$ (Ponomarev et al. 1970). The experimentally known phases in which aluminum atoms in mixed coordination have been described: $\text{Ca}_2\text{Al}_2\text{O}_5$ (Kahlenberg et al. 2000b; Lazić et al. 2008) with aluminum in tetrahedra and octahedra; $\text{Ca}_4\text{Al}_6\text{O}_{13}$ (Kahlenberg et al. 2000a), with aluminum in tetrahedra, octahedra and trigonal bipyramids; $\text{Ca}_3\text{Al}_2\text{O}_6$ (Steele and Davey 1929) with Al in octahedra and squares, and $\text{CaAl}_{12}\text{O}_{19}$ (Kato and Saalfeld 1968) with aluminum in octahedra and trigonal bipyramids. The mineral mayenite (Hentschel G., 1964), which has a clathrate structure, has a rather complex composition $\text{Ca}_{12}\text{Al}_{14}\text{O}_{33}$ and the removal of one oxygen atom per formula unit leads to a peculiar metallic electroneutral state (Matsuishi et al. 2003) (Figure S1). Overall, the tetrahedral coordination of aluminum is unstable at pressures above 2–8 GPa (Ito et al. 1980; Irifune and Tsuchiya, 2007; Eremin et al. 2016), therefore, phases with tetrahedral coordination of Al atoms in the $\text{CaO--Al}_2\text{O}_3$ system cannot exist in the deep mantle.

In this work, we theoretically predicted thermodynamically stable phases in the $\text{CaO--Al}_2\text{O}_3$ system at pressures in the range of 0–200 GPa using the Universal Structure Predictor: Evolutionary Xtallography (USPEX) algorithm (Oganov and Glass 2006; Oganov et al. 2011; Lyakhov et al. 2013).

Methods

Evolutionary searches for stable phases in the CaO-Al₂O₃ system were performed using the USPEX code (Oganov and Glass 2006; Oganov et al. 2011; Lyakhov et al. 2013) in its variable-composition mode. In our evolutionary searches, all structures and stoichiometries were allowed with up to 32 atoms in the primitive cell. The first generation contained 200 structures and was produced using symmetric (Lyakhov et al. 2013) and topological (Bushlanov et al. 2019) random structure generators. All subsequent generations had 120 structures. To determine stable compounds, USPEX builds convex hull diagrams (also known as Maxwell construction). Points lying on the hull indicate thermodynamically stable phases. Points lying above the hull are thermodynamically unstable phases. Evolutionary searches were performed at pressures of 25, 50, 100, and 200 GPa. All structures were relaxed using VASP (Kresse and Furthmüller 1996), using density functional theory at the level of generalized gradient approximation, with PBE exchange–correlation functional (Perdew et al. 1996) and PAW method (Blöchl 1994; Kresse and Joubert, 1999), with plane-wave cutoff of 600 eV. Brillouin zone was sampled using uniform k-meshes with reciprocal-space resolution $2\pi \times 0.05 \text{ \AA}^{-1}$, which ensures that the errors in energy differences are below 1 meV/atom. Low-enthalpy structures predicted using USPEX and experimentally known structures of different stoichiometries were carefully relaxed at pressures ranging from 0 to 200 GPa. The full zero-temperature pressure-composition phase diagram of the CaO-Al₂O₃ system was constructed. It is worth noting that taking into account thermal contributions to the Gibbs free energy requires a lot of computing resources and is not required in this study, because the entropic factor ($T\Delta S$) at the temperatures of the Earth's mantle is estimated to be much less than $k_B T$ (where k_B is the Boltzmann constant). Thus, the entropy contribution under such P - T conditions is negligibly small compared to ΔH and was not taken into account.

Thus, the phase stability is unlikely to be changed qualitatively by temperature at pressures of the lower mantle. Crystal structures were visualized using the VESTA program (Momma and Izumi 2011).

Results and discussion

Variable-composition USPEX searches allow for a detailed sweep of the entire compositional space in a single simulation. They recovered all known stable structures and phase transitions between them in CaO (Jerebtsov and Mikhailov

2001; Deng et al. 2007) and Al₂O₃ (Marton and Cohen 1994; Oganov and Ono 2005).

Interestingly, USPEX has also found (as stable or metastable at zero Kelvin) all previously known structures of CaAl₂O₄ with the octahedral coordination of aluminum atoms, including the marokite-type phase and the "layered" $P2_1/m$ modification. As we find, the $Pnma$ modification of CaAl₂O₄ with a marokite structure type turns out to be stable at pressures from 7.0 to at least 200 GPa (Figs. 2 and 3). It is worth noting that the previously theoretically predicted (Eremin et al. 2016) metastable CaAl₂O₄ phase with α -PbO₂-like structure is strongly inferior in energy to marokite-type phase in the studied pressure range.

At pressures below 7.0 GPa, the $P2_1/m$ structure is stable, and below 1.1 GPa the $P2_1/n$ structure is stable (Fig. 2). The $P2_1/c$ phase is unstable over the entire pressure range. At low pressures, we see yet another two phases as stable: 5CaO • 3Al₂O₃ (C_5A_3) with space group $Cmc2_1$ (up to 2.1 GPa) and CaAl₄O₇ with space group $C2/c$ (up to 1.8 GPa) (Fig. 3a). However, the stability of C_5A_3 phase is marginal: if we include zero-point energy (we did its crude estimate with Debye model), C_5A_3 phase is no longer thermodynamically stable but is 0.009 eV/block above the convex hull. The convex hull of the CaO-Al₂O₃ system at pressures of 0 GPa and 100 GPa are shown in Fig. 4. It should be noted that the Ca₃Al₂O₆ (C_3A) and CaAl₁₂O₁₉ (CA_6) and Ca₁₂Al₁₄O₃₃ ($C_{12}A_7$) phases at low pressures are very close to the convex hull, being just 0.008, 0.027, and 0.008 eV/block, respectively, above it at a pressure of 0 GPa. The calculated fitness (the height above the convex hull, characterizing the degree of instability of a given structure) for these experimental structures as a function of pressure is shown in Figure S3. The calculated enthalpy of different polymorphs with CA composition is shown in Table S2.

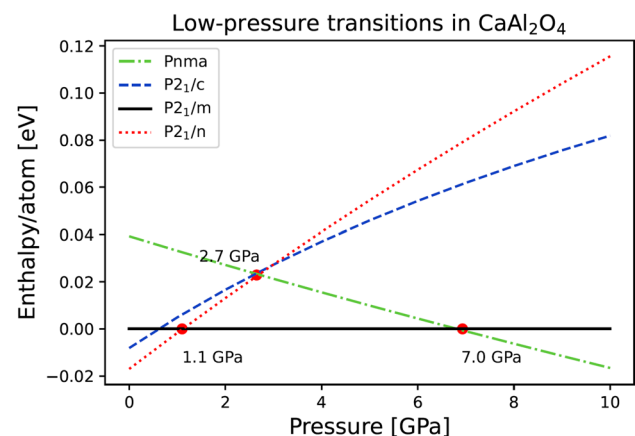


Fig. 2 Enthalpy differences between low-pressure phases of CaAl₂O₄. Enthalpy for $P2_1/m$ phase was taken as a reference

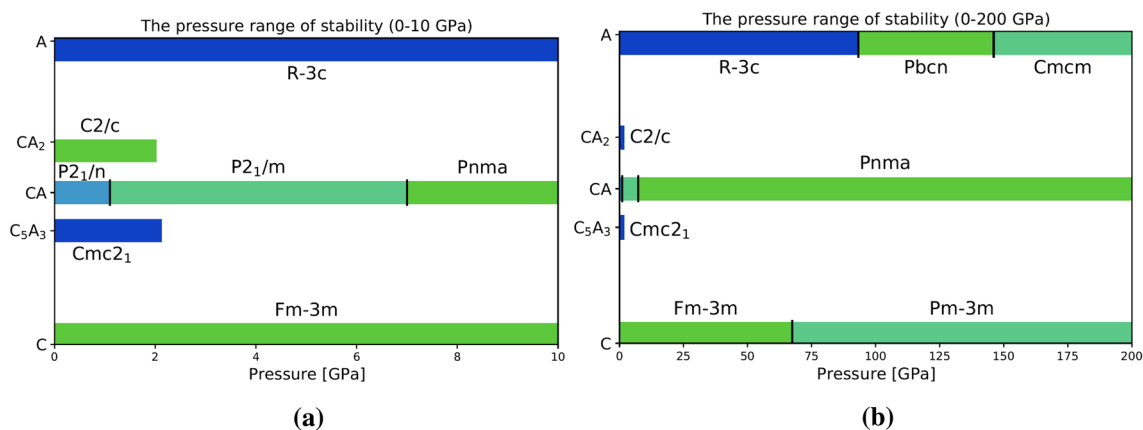


Fig. 3 Pressure-composition phase diagram of the CaO-Al₂O₃ system **a** in the 0–10 GPa pressure range, **b** in the 0–200 GPa pressure range

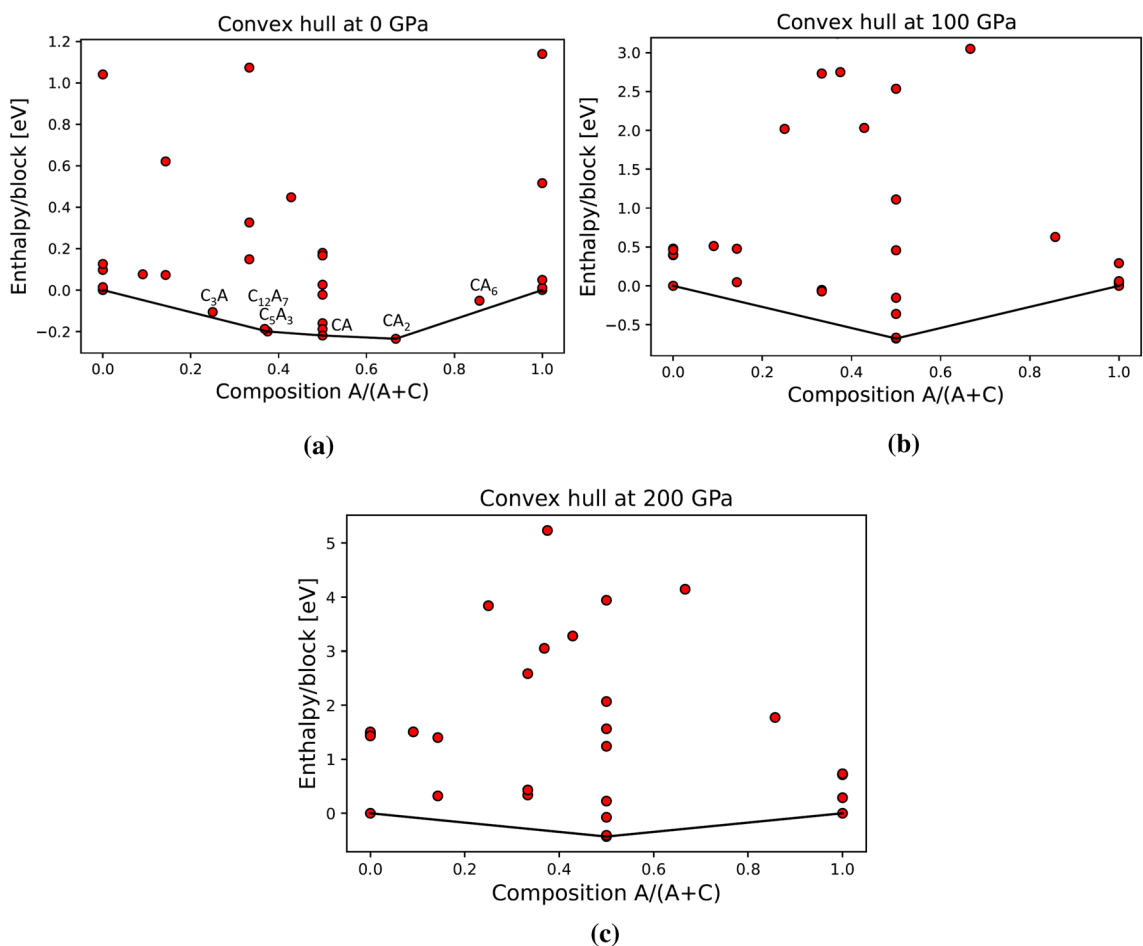


Fig. 4 Convex hulls of the CaO(C)-Al₂O₃ (A) system at 0 GPa **a**, 100 GPa **b**, and 200 GPa **c**

We also calculated phase transition pressures for CaO and Al₂O₃. For CaO, the pressure of the phase transition from *Fm* $\bar{3}m$ to *Pm* $\bar{3}m$ is 67.4 GPa (Figure S2a). For Al₂O₃, the pressure of phase transitions from *R* $\bar{3}c$ to

Pbcn is 93.2 GPa and further from *Pbcn* to *Cmcm* 146.1 GPa (Figure S2b). This compares well with the experiments (Jeanloz et al. 1979)(Funamori and Jeanloz 1997)

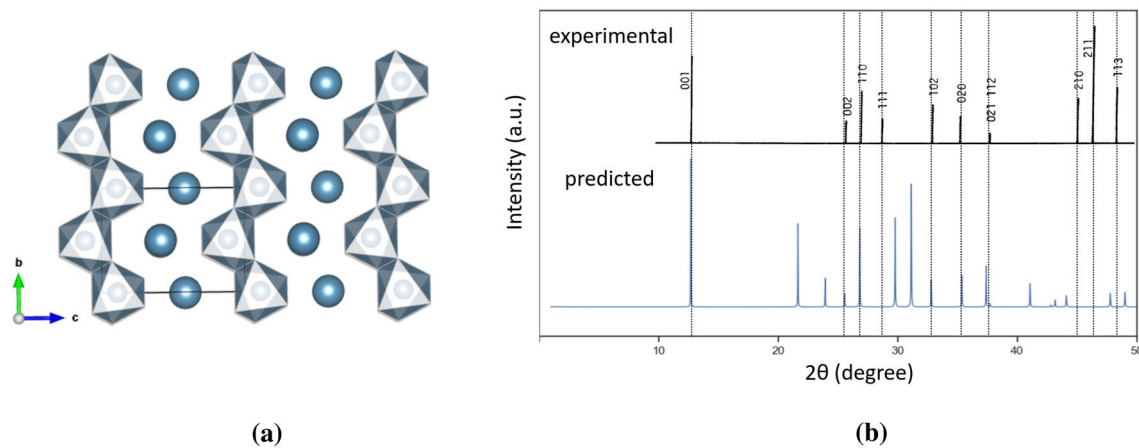


Fig. 5 Predicted structure of "CA-III" metastable phase **a** and diffraction patterns of the "CA-III" phase according to the experimental data from Ito et al. (1980) and our predicted structure of "CA-III" phase **b**. The AlO_6 polyhedra are shown in gray color, Ca^{2+} ions are shown as large balls

and previous theoretical calculations (Oganov and Ono 2005; Deng et al. 2007; Liu et al. 2015).

To find the unknown crystal structure of the "CA-III" phase (Ito et al. 1980) with cell parameters $a = 4.390 \text{ \AA}$, $b = 5.070 \text{ \AA}$, $c = 6.960 \text{ \AA}$, we performed evolutionary USPEX searches with a fixed CaAl_2O_4 and fixed experimental cell parameters. The resulting structure has indeed the $P2_12_12$ symmetry, and features layers made of edge- and face-sharing AlO_6 octahedra (Fig. 5a), with Ca atoms located in the interlayer space. The coincidence of the main X-ray diffraction peaks ((001), (002), (110), (102), (020) and (112)) of our structural model and the experimental (Ito et al. 1980) diffraction patterns (see Fig. 5b) allows us to conclude that the structure was identified correctly. It should be noted that texturing of the sample in the experiment leads to a difference between the theoretical and experimental diffraction peak intensity ratios. The predicted crystal structure of the "CA-III" phase is metastable at 0 K, being 0.18 eV/atom higher in enthalpy than the monoclinic $P2_1/n$ phase.

Conclusions

Our systematic evolutionary searches for stable compounds in the $\text{CaO-Al}_2\text{O}_3$ system at pressures 0–200 GPa clearly establish that at pressures above 7.0 GPa (i.e. at all pressures of the Earth's transition zone and lower mantle) only one calcium aluminate, CaAl_2O_4 with CaFe_2O_4 structure ($Pnma$), is stable. This phase is one of the prime candidate aluminous phases in the lower mantle of the Earth. En route to this conclusion, we have also resolved the old puzzle of the low-pressure (3–8 GPa) metastable phase named "CA-III" – its crystal structure is now established as having the orthorhombic space group $P2_12_12$, and featuring layers of

edge- and face-sharing AlO_6 octahedra. Our analysis also reveals that at low pressures $5\text{CaO} \bullet 3\text{Al}_2\text{O}_3$ (C_5A_3) with space group $Cmc2_1$, CaAl_4O_7 ($C2/c$) and CaAl_2O_4 ($P2_1/m$) structures are stable up to 2.1, 1.8 and 7.0 GPa respectively.

CaAl_2O_6 , $\text{CaAl}_{12}\text{O}_{19}$ and $\text{Ca}_{12}\text{Al}_{14}\text{O}_{33}$ phases are very close to thermodynamic stability at low pressures and are also known from the experiment. Very rich and complex at low pressures, the chemistry of the cement $\text{CaO-Al}_2\text{O}_3$ system becomes very simple at high pressures.

Competing interests

The authors declare that they have no competing interests.

Supplementary Information The online version contains supplementary material available at <https://doi.org/10.1007/s00269-022-01221-6>.

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Author contributions EI. M: conceptualization, investigation, writing—original draft preparation; AR. O: data curation, writing—reviewing and editing; EA. M: investigation, methodology; NN. E: conceptualization, writing—reviewing and editing.

Declarations

Competing interests The authors declare no competing interests.

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